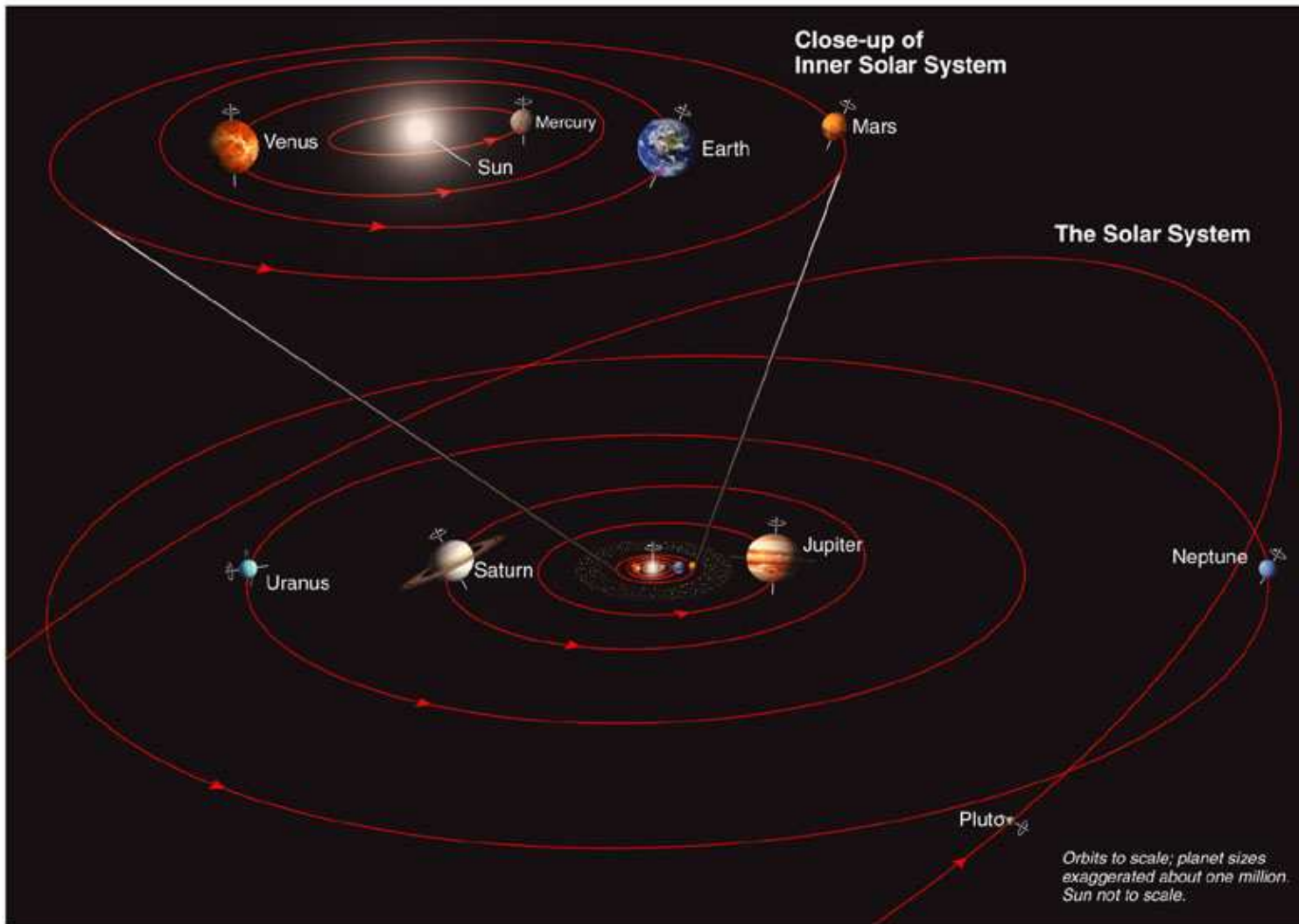


Solar System overview

- 1) inventory
- 2) spin/orbit/shape
- 3) heated by the Sun
- 4) how do we find out



Inventory

1 star

(99.9% of M)

8 planets

(99.9% of L)

- Terrestrial:

Mercury

Venus

Earth

Mars

- Giant:

Jupiter

Saturn

Uranus

Neptune

Lots of small bodies

incl. dwarf planets

Ceres

Pluto

Eris

Inventory (cont'd)

Many moons & rings

Mercury: 0

Venus: 0

Earth: 1 (1700km)

Mars: 2 (~10km)

Jupiter: 63 + rings

Saturn: 60 + rings

Uranus: 27 + rings

Neptune: 13 + rings

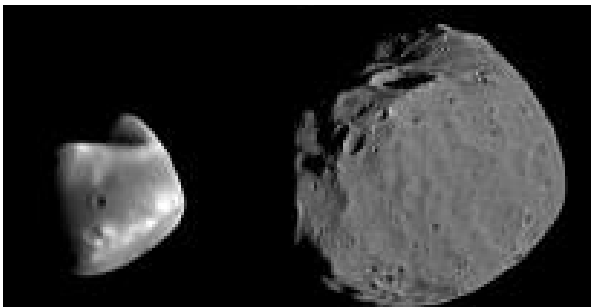
Even among dwarf planets, asteroids, Kuiper belt objects, and comets. E.g.,

Pluto: 3

Eris: 1

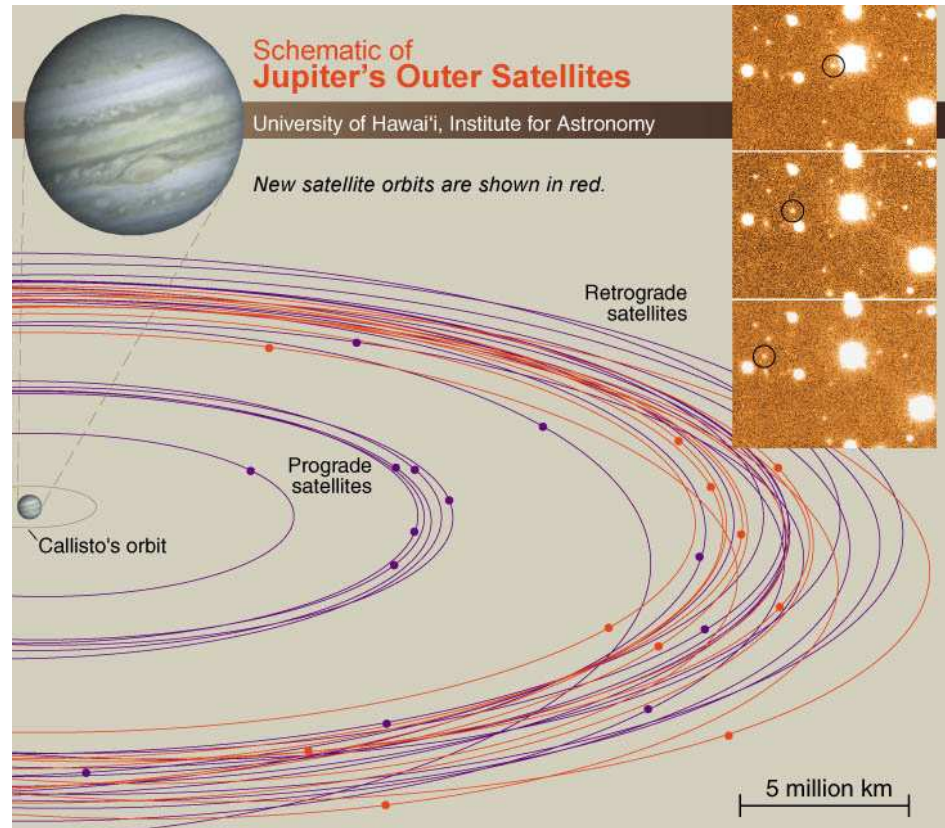
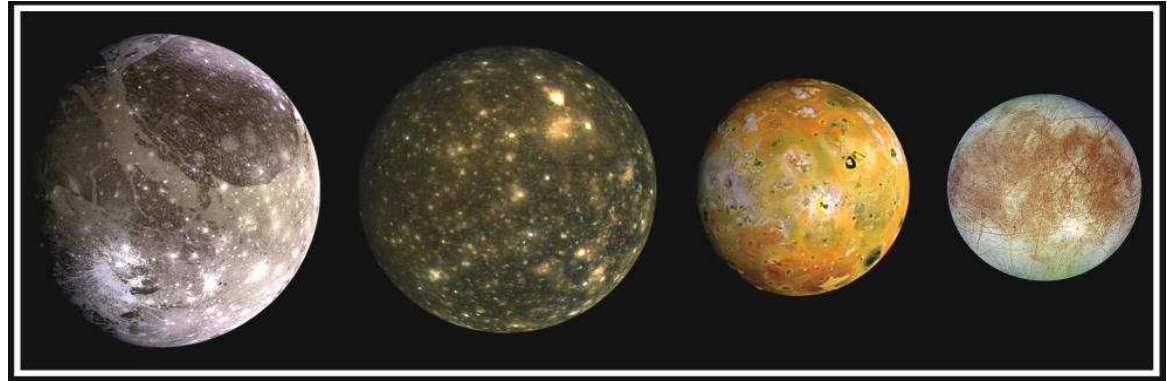
Moons of Mars:

Deimos & Phobos, ~10km



Moons of Jupiter

4 Galilean satellites (Ganymede, Callisto, Io & Europa),
~ 10^3 km (close to Jupiter, likely primordial)



2001J3: 4km

Atmosphere

no



MERCURY
4,878 KM
(3,024 MI)

thick



VENUS
12,100 KM
(7,502 MI)

thick



EARTH
12,756 KM
(7,909 MI)

little



MARS
6,798 KM
(4,214 MI)

thick



MOON (EARTH)
3,476 KM
(2,155 MI)



IO (JUPITER)
3,630 KM
(2,251 MI)



EUROPA (JUPITER)
3,138 KM
(1,946 MI)



GANYMEDE (JUPITER)
5,262 KM
(3,262 MI)



CALLISTO (JUPITER)
4,800 KM
(2,976 MI)



TITAN (SATURN)
5,150 KM
(3,193 MI)

Inventory (cont'd)

~ 10^5 known small objects in the

- **Asteroid belt**
(Ceres ~300 km)

- **Kuiper belt**
(Eris, Pluto, Sedna, Quaoar, ~1000 km)

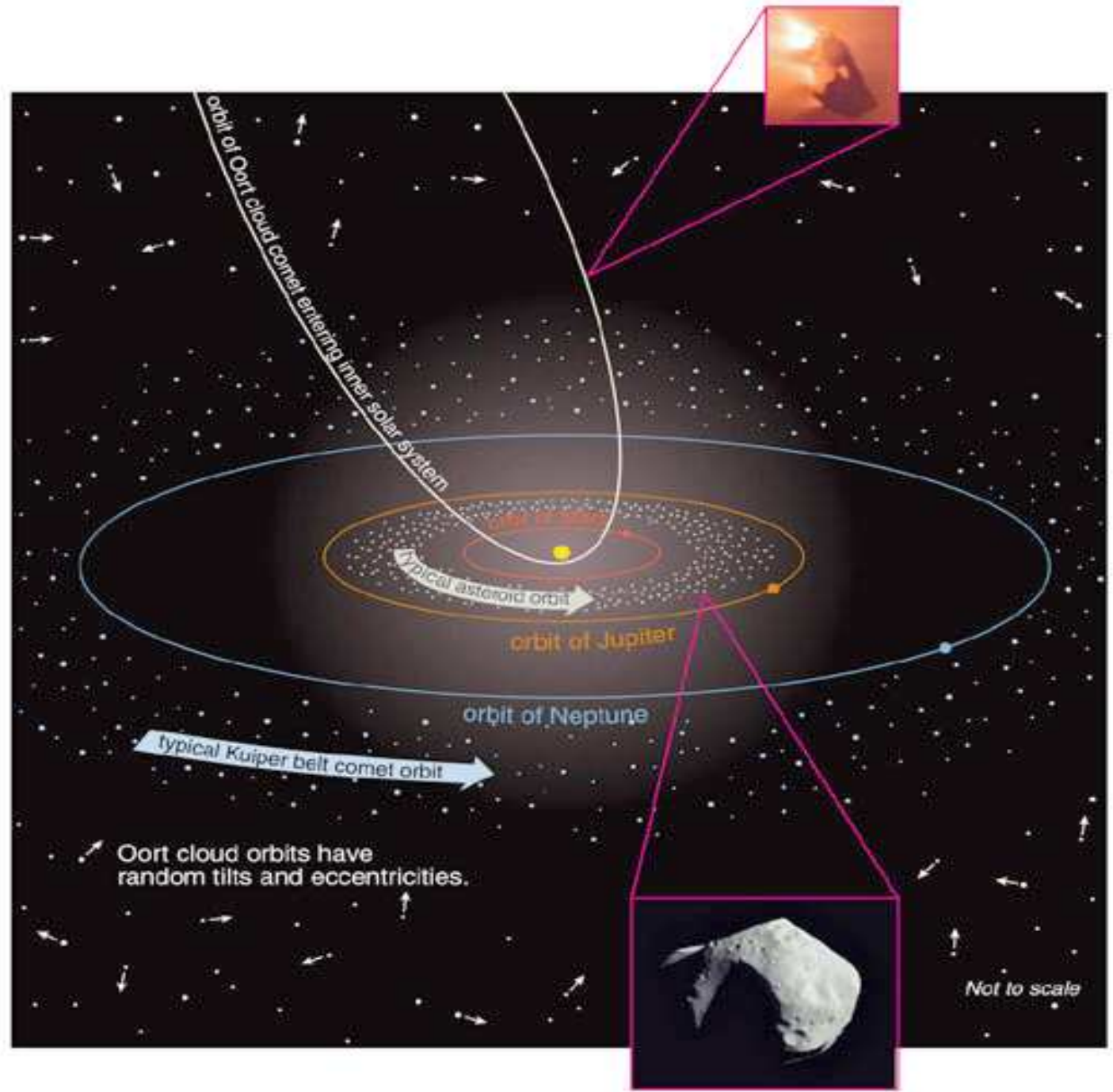
Estimated: ~ 10^{12} comets in the

- **Oort cloud**
(~ 10^4 AU)

Associated:

- **zodiacal dust**

(fire-works on the sky:
comets & meteorites)



What are planets?

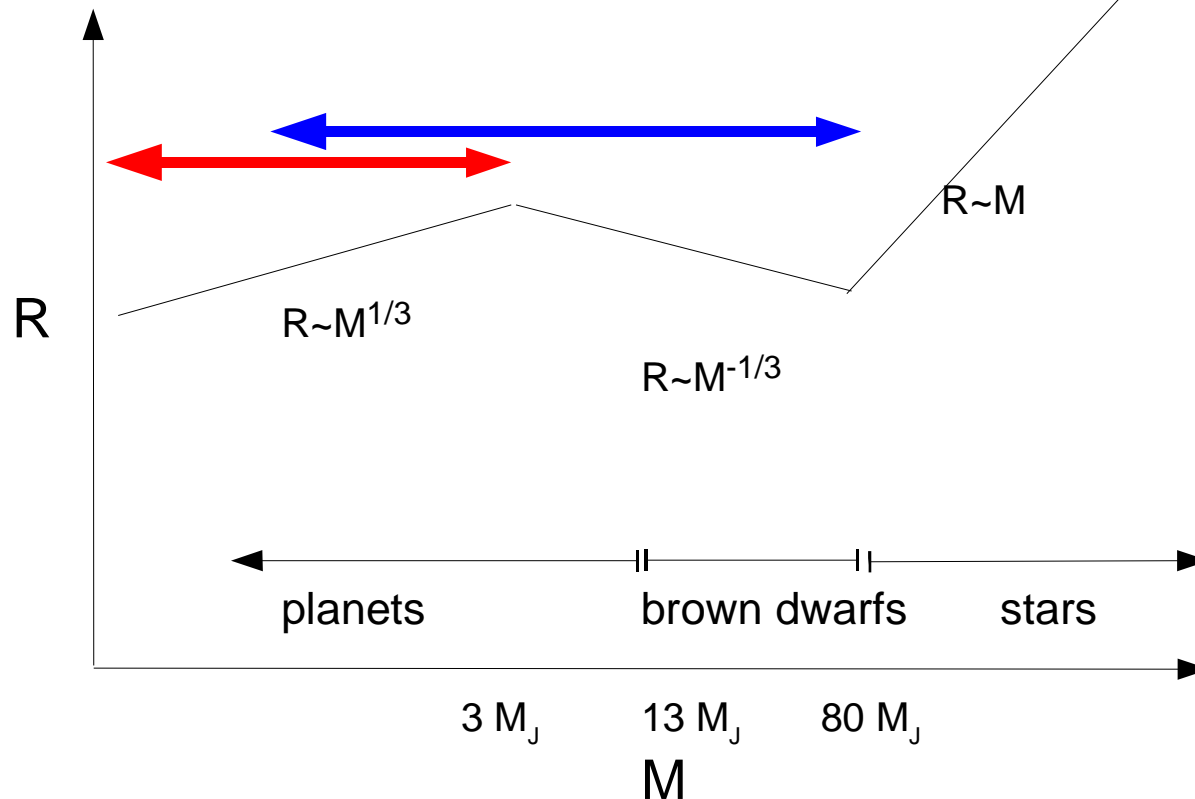
IAU (for solar system):

Orbits Sun, massive enough to be round and to have cleared its neighbourhood.

More general:

- 1) no nuclear fusion (not even deuterium): $T_c < 10^6 \text{ K}$
- 2) pressure provided by **electron degeneracy** and/or **Coulomb force**
 $(l \sim h/p \sim d)$ $(d \sim \text{atomic radius})$

- 3) can be solid or gaseous (with solid cores) --- *similar density*



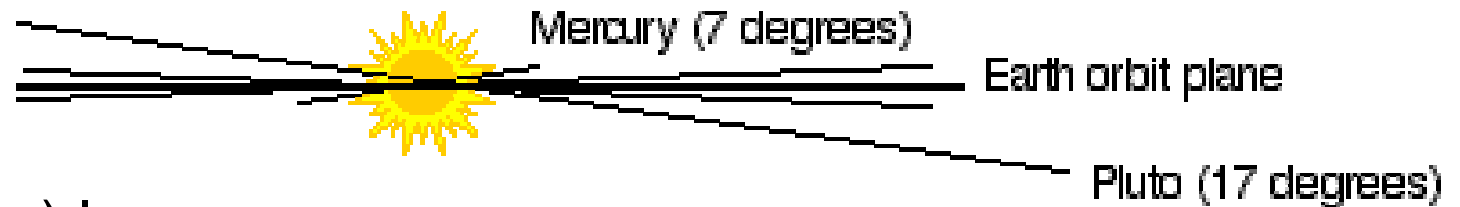
Mass & Mean ρ		
	M_J	[g/cm ³]
Jupiter	1.0	1.33
Saturn	0.3	0.77
Neptune	0.05	1.67
Uranus	0.04	1.24
Earth	0.003	5.52
Venus	0.002	5.25
Mars	0.0003	3.93
Mercury	0.0002	5.43

Orbits

inclination: largely coplanar (history)

direction: all the same

eccentricity: a few percent (except for Mercury)



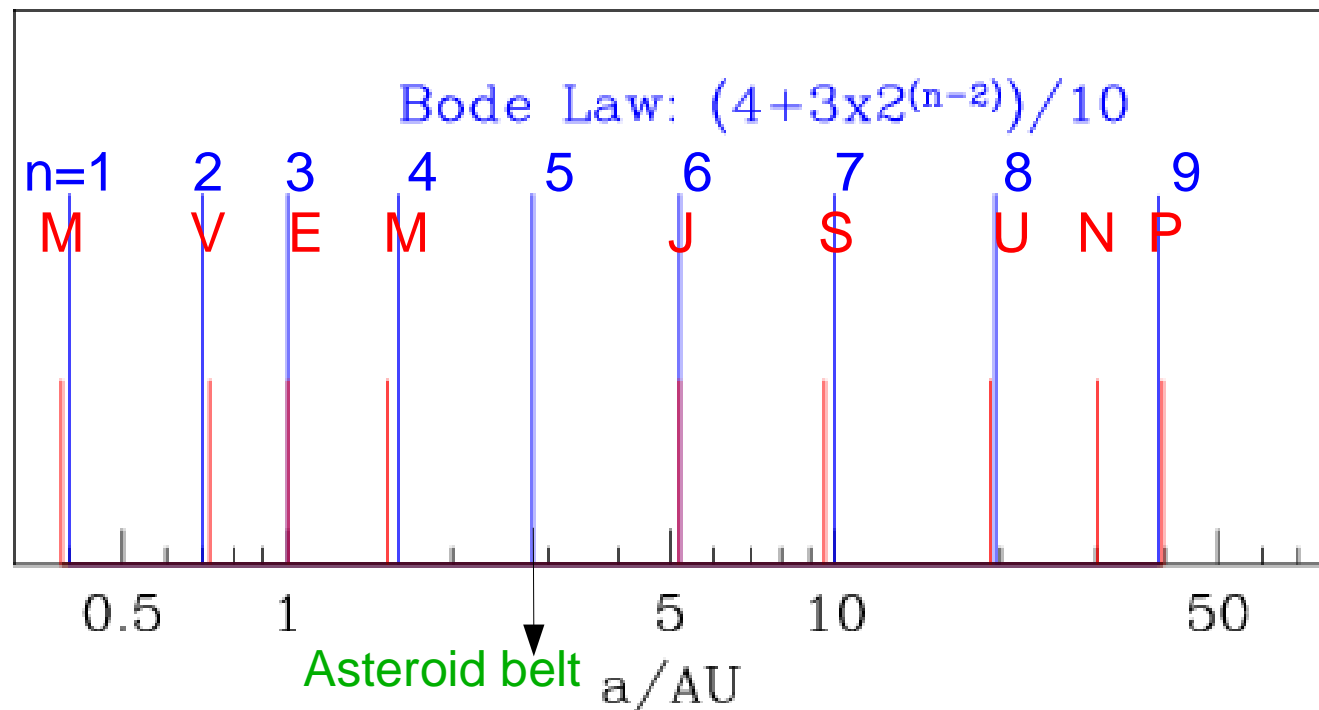
Titus-Bode (fitting) law (1766)

planetary orbits appear to (almost) satisfy a single relation

'Predict' the existence of the asteroid belt (1801: Ceres discovered)

coincidence or something deeper?

other systems?

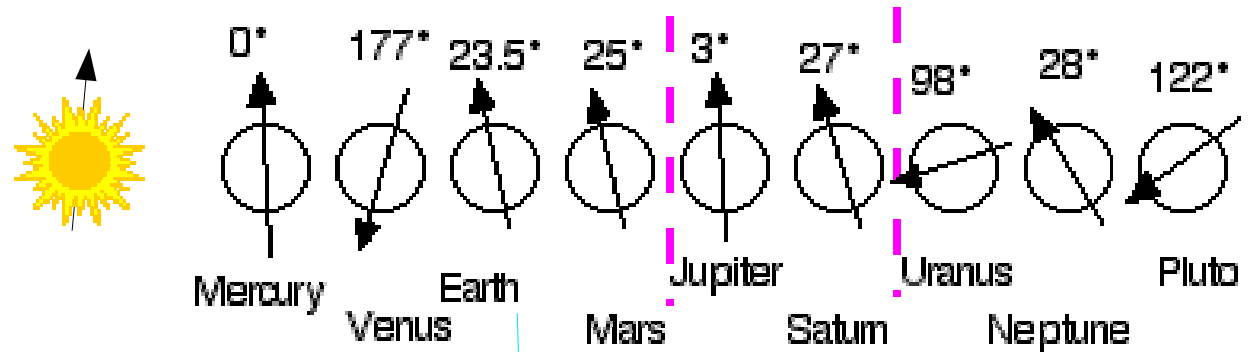


computer simulations indicate that planets are as maximally packed as allowed by stability

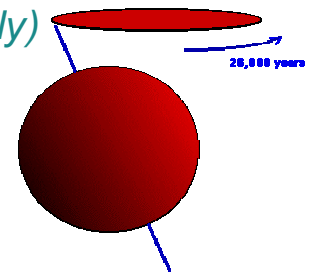
Spin (obliquity)

smaller planets:
almost random, affected by
impacts and giant planets

Real giant planets (J&S):
~aligned with orbit, stable



*Earth's spin-axis precesses (mildly)
while Mars sweeps around wildly*



Shape --- the bigger the rounder

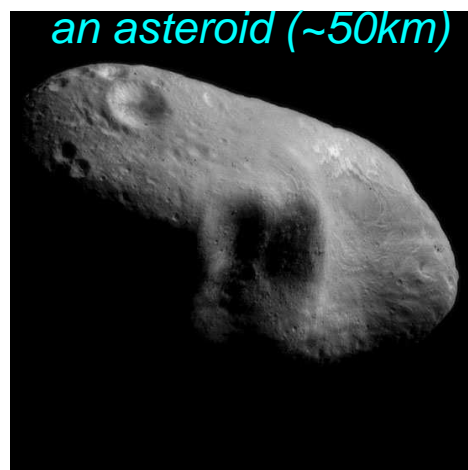
All gaseous planets are spherical.

Large rocky objects are rather spherical. Smaller ones are less so.

The Moon (~1700km)



an asteroid (~50km)



	h	R	$g=GM/R^2$
	(km)	(km)	(m/s ²)
Earth	8	6400	9.8
Mars	24	3400	3.7

scaling: highest mountain on Earth ~8 km (on Mars ~ 24 km) **$h * g \sim \text{constant}$**

rough estimate: irregular body has mountain $h \sim R \implies R \sim 240 \text{ km}$

thus: objects with $R > 240 \text{ km}$ are approximately spherical

Saturn's Eight Major Icy Satellites

400 km



Mimas



Enceladus



Tethys



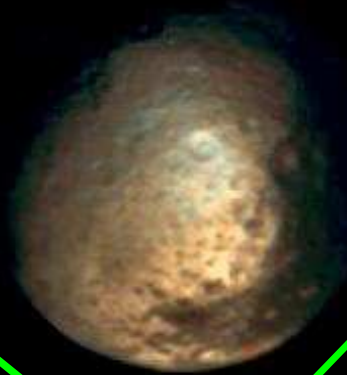
Dione



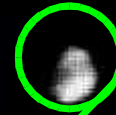
Rhea



Hyperion

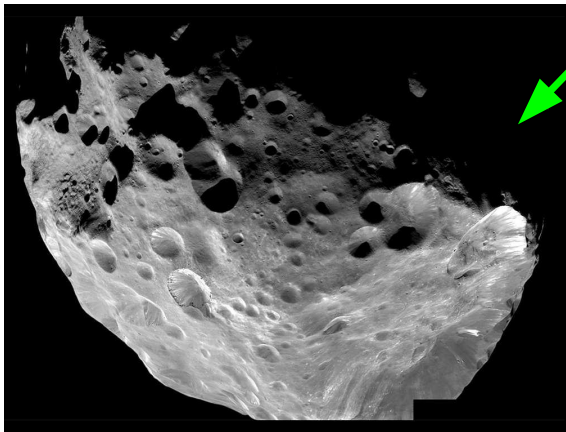


Iapetus



Phoebe

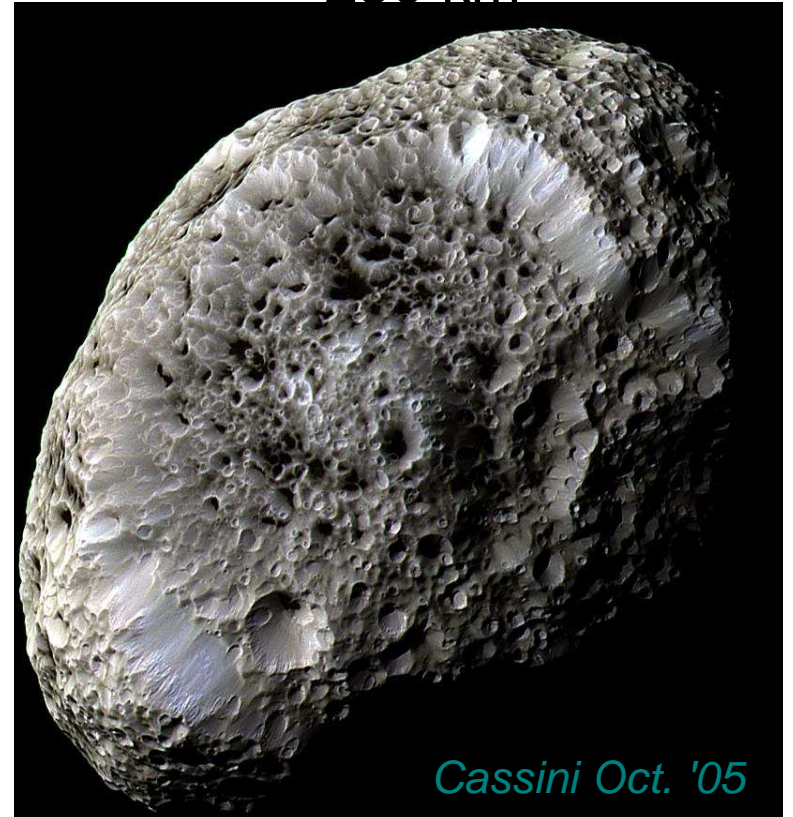
~200 km



The bigger the rounder

	$\Delta R/R$	$g=GM/R^2$
Earth	8/6400	9.8 m/s ²
Mars	24/3400	3.7 m/s ²
Hyperion: 150/250		~0.4 m/s ²

~250 km



Cassini Oct. '05

Passively Heated by the Sun --- *the further the cooler*

Typically we observe objects in reflected light, however, all objects emit re-processed thermal radiation which is observable at longer wavelengths.

Blackbody temperature for a non-self-luminous spherical body at distance a away from the Sun (with albedo A -- reflectivity)

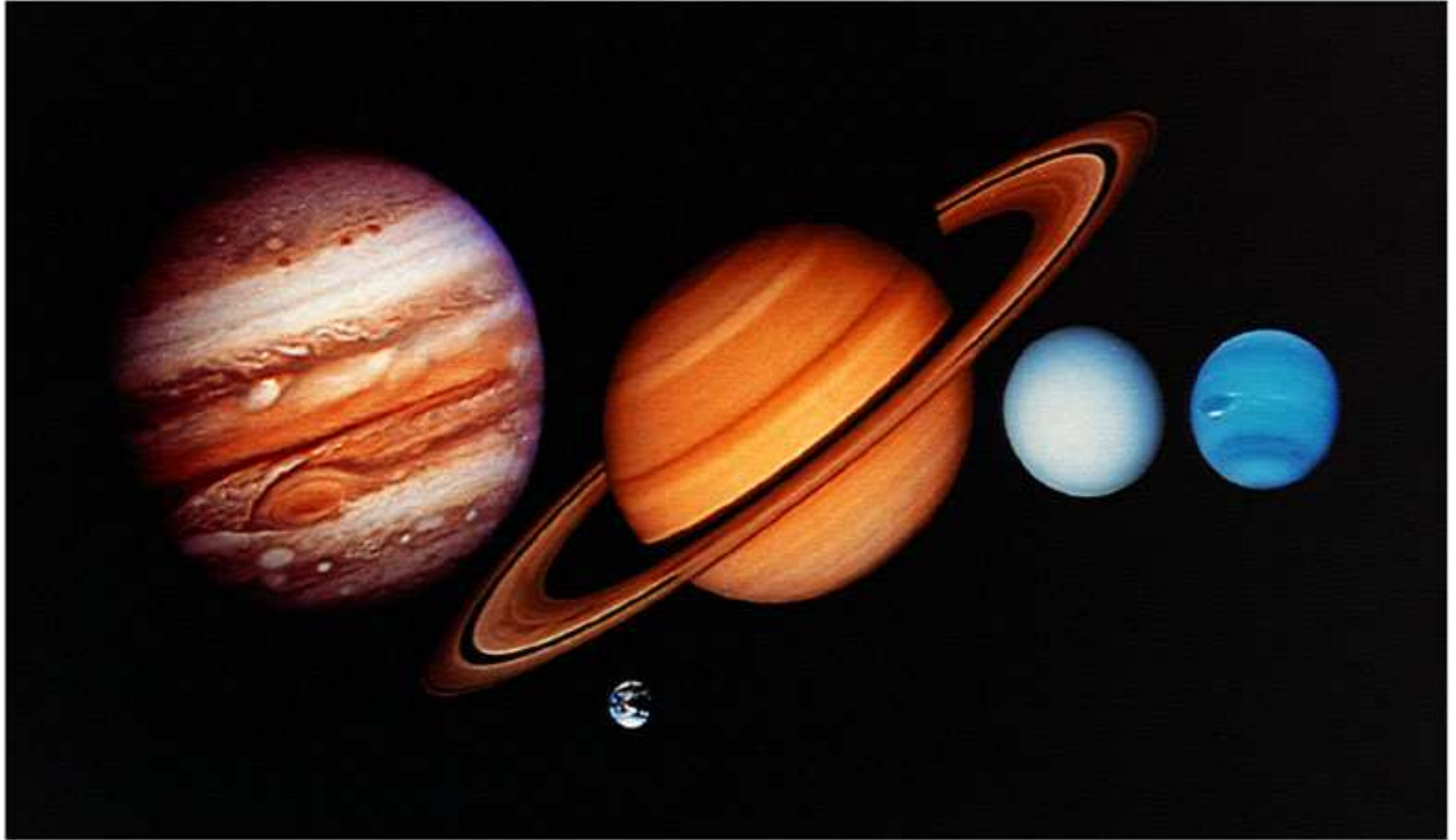
$$L_{\text{abs}} = (1-A) \frac{\pi R_p^2}{4\pi a^2} 4\pi R_s^2 \sigma T_s^4; \quad L_{\text{em}} = 4\pi R_p^2 \sigma T_p^4$$

If $L_{\text{abs}} = L_{\text{em}}$, then

$$T_p = \left(\frac{R_o}{2a} \right)^{1/2} T_s (1-A)^{1/4}$$

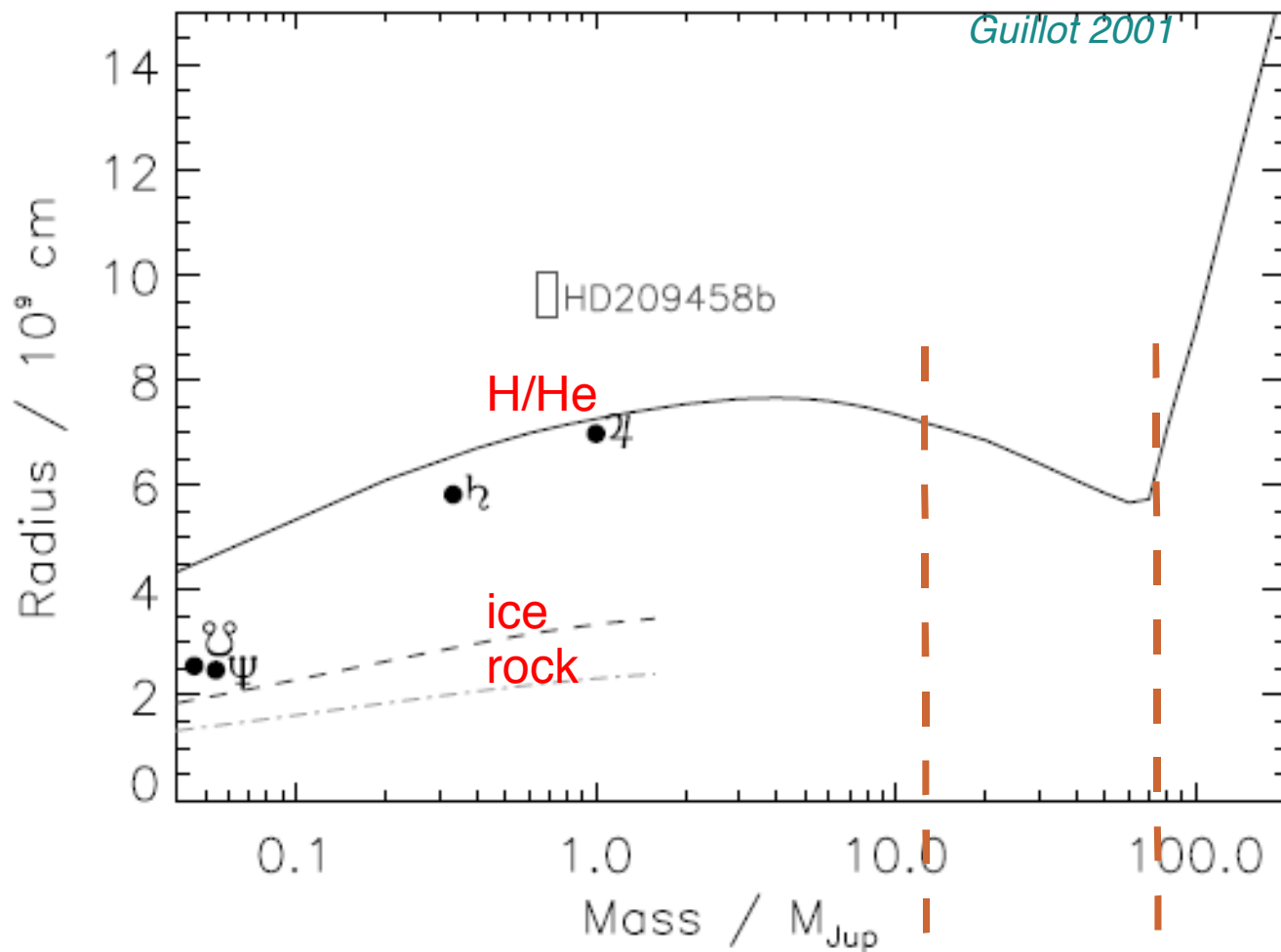
	a (AU)	A	$T_{\text{pred}}(\text{K})$	$T_{\text{act}}(\text{K})$	
Mercury	0.4	0.06	422 K	100-725	(?)
Venus	0.7	0.77	230K	733	(?)
Earth	1	0.30	255K	288	(?)
Mars	1.5	0.25	218K	223	good
Jupiter	5	0.51	113K	125	(?)
Saturn	9	0.47	83K	95	(?)
Uranus	19	0.51	60K	60	good
Neptune	30	0.62	40K	60	(?)
Comet at	5000	0.51	3.4K		

Giant Planets



made mostly of H, He and H-compounds, no solid surface
99.5% planet mass, 99.8% solar system angular momentum

Giant planets border stars



Coulomb pressure
planets

e- deg. pressure
brown dwarfs

ideal gas
stars

13 M_J

80 M_J

Equation of state determines mass-radius relation

Ideal gas: $P \propto \rho T$
 +Fusion: $T_{core} \sim \text{Const}$
 $\rightarrow R \propto M$

el. degeneracy: $P \propto \rho^{5/3}$
 $\rightarrow R \propto M^{-1/3}$

Coulomb: $\rho \sim \text{Const}$
 $\rightarrow R \propto M^{1/3}$

Working definition:
Brown-dwarfs are 'failed' stars that cannot ignite hydrogen (but can burn deuterium); hence $M < 80 M_J$ ($0.08 M_{\odot}$)

Planets are formed in disks around stars. Planets cannot burn deuterium (10^6 K); hence $M < 13 M_J$

Are planets just gas balls like stars? Probably not.

Jupiter & Saturn: largely degenerate H & He, mean $\rho = 1.3$ & 0.7 g/cm^3

-- hydrogen metallic (conductive) below certain depth (?)

-- core: solid, heavy metal + ices

Jupiter's core: $< 10 M_E$ (or 0?); Saturn's core: $\sim 13 M_E$ (15% of mass)

Uranus & Neptune: largely ices (H_2O , CH_4 , NH_3), mean $\rho = 1.2$ & 1.7 g/cm^3

-- relatively thin gaseous H & He envelope

-- mostly icy + rocky core

Why do we care about the solid cores?

Formation of giant planets likely starts with a solid core – unlike stars

How do we figure out about the cores? Spin it!

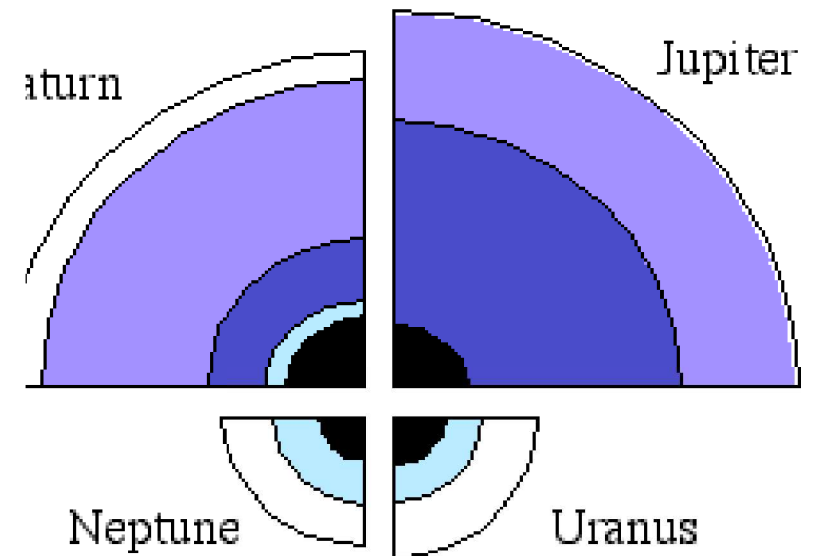
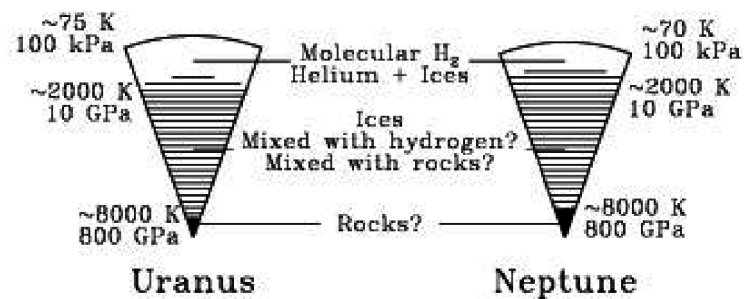
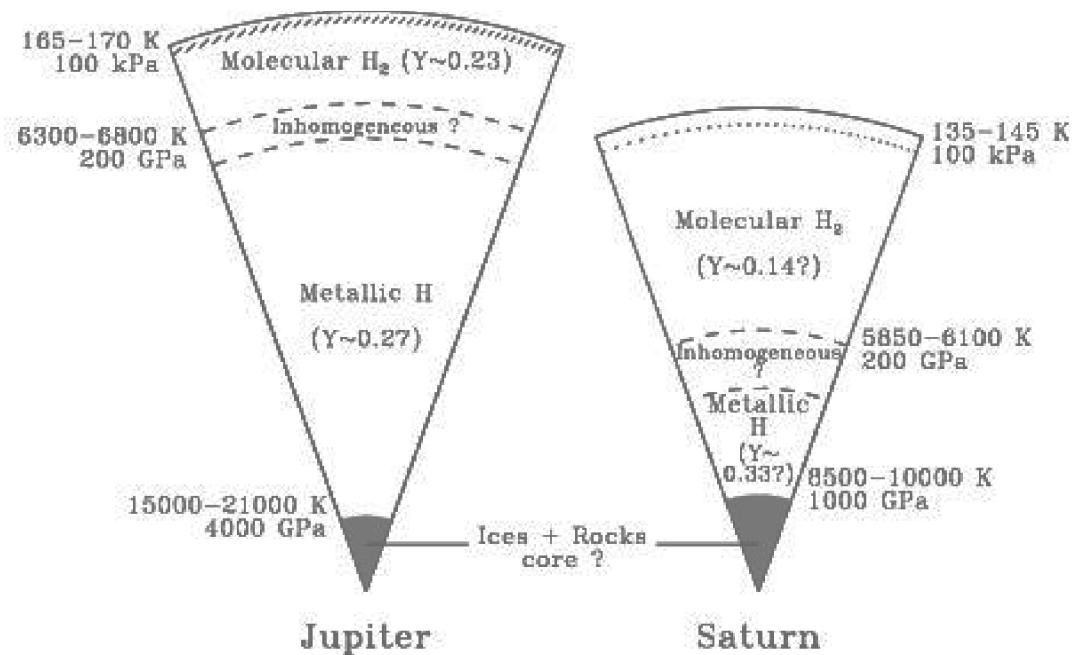
core: a high density central region

spherical body: gravitational potential is independent of density profile

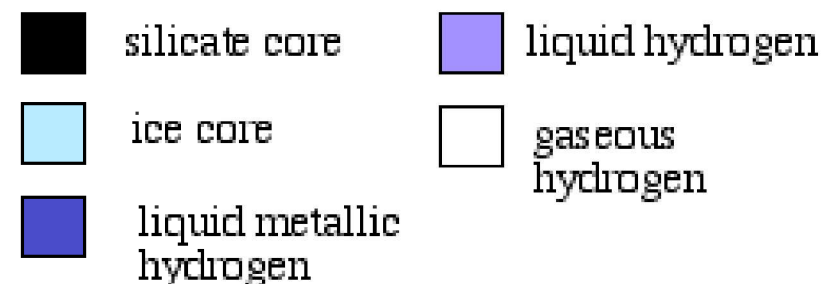
but when the planet rotates, its oblateness depends on $\rho = \rho(r)$

$$\Phi(\theta) = -\frac{GM}{r} \left[1 - \left(\frac{R}{r} \right)^2 J_2 P_2(\cos \theta) - \left(\frac{R}{r} \right)^4 J_4 P_4(\cos \theta) - \dots \right]$$





● Earth to same scale



Cores of giant planet are likely primordial and do not form by gravitational settling.
Did Jupiter melt part of its core?

Energy budget for giant planets

$$\text{Absorb solar flux: } (1 - A) 4 \pi R_o^2 \sigma T_o^4 \times \frac{\pi R_p^2}{4 \pi a^2}$$

$$\text{Emit blackbody flux: } 4 \pi R_p^2 \sigma T_p^4$$

$$T_p = (1 - A)^{1/4} \left(\frac{R_o}{2a} \right)^{1/2} T_o$$

	Jupiter	Saturn	Uranus	Neptune
passive T_p	113K	83K	60K	48K
actual T_p	130K	95K	59K	59K
$L_{\text{total}}/L_{\text{received}}$	1.7	1.8	1.0	2.6

3 sources of planetary intrinsic luminosity: primordial + settling + radio-active

Jupiter: **primordial heat** + He settling relative to H
(very long thermal time-scale: $\sim 10^9$ yrs)

Saturn: primordial heat + He settling relative to H

Uranus: no additional source required

Neptune: Do require add'l source; but so similar to Uranus, so why?

--- *what about gravitational contraction?* No, already shrunk

--- *terrestrial planets: radio-active elements*

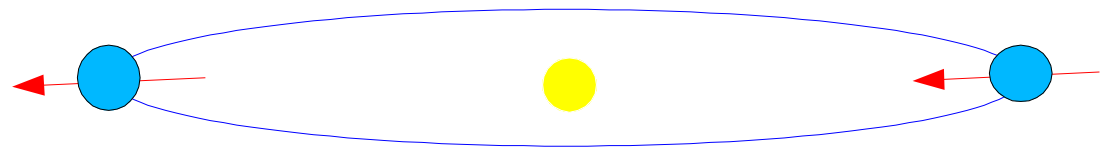
--- *how much energy can you gain by separating H & He?*

Other cool points?

1) magnetic fields: all 4 have appreciable B fields, Jovian aurorae,
Jupiter's magnetic influence extends past Saturn orbit
generation of these fields -- primordial or dynamo?

2) seasons:

Uranus: 97.92° inclined relative to orbit, very weird seasons!



3) rings & satellites: all 4 have rings and many satellites

rings: sandy or icy dust and some boulders, 2.5 planet radii (\sim Roche radius)

-- $H/R \sim 10^{-6}$ (*a razor blade?*)

--- gaps: shepherding moons

-- origin: tidally disrupted satellites
or primordial?

Satellites: worlds of their own

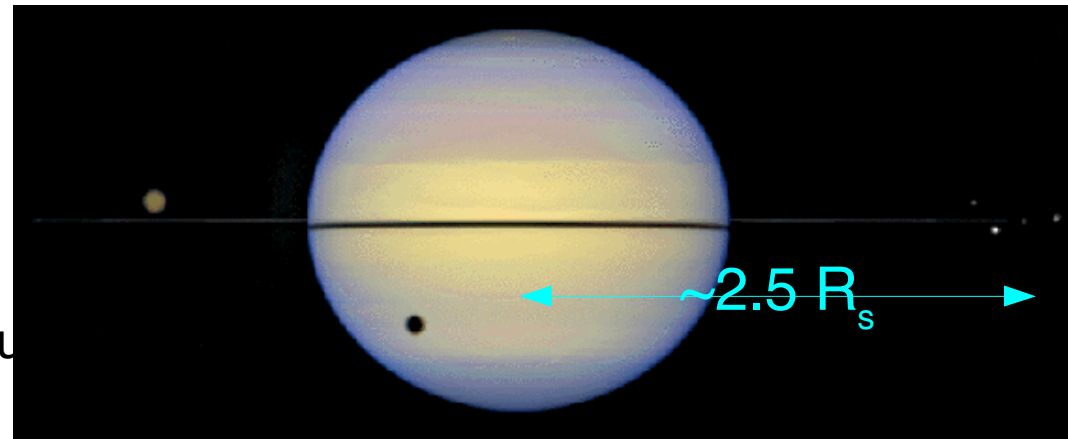
captured (Phoebe) or formed in-situ

Europa (@J): cracky surface

underground H_2O ocean

Titan (@S): smoggy atmosphere

surface H-compound ocean?



Saturn's rings (Cassini images)

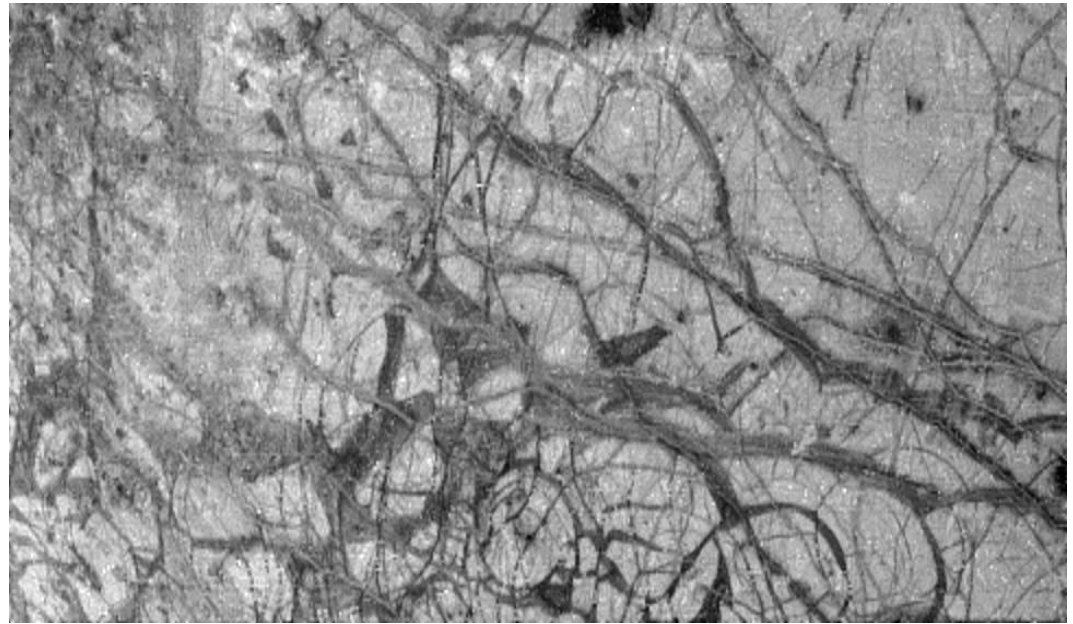
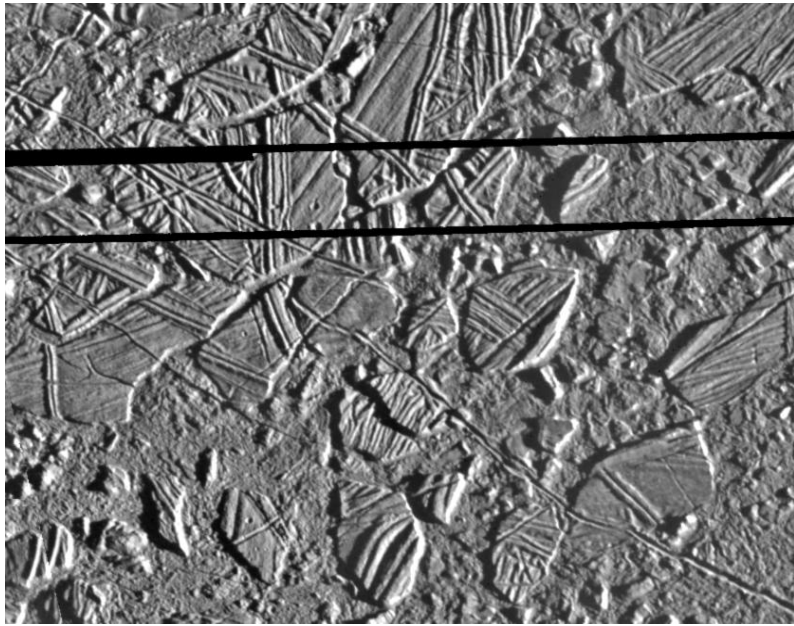
Sharp edges

Prometheus shepherding

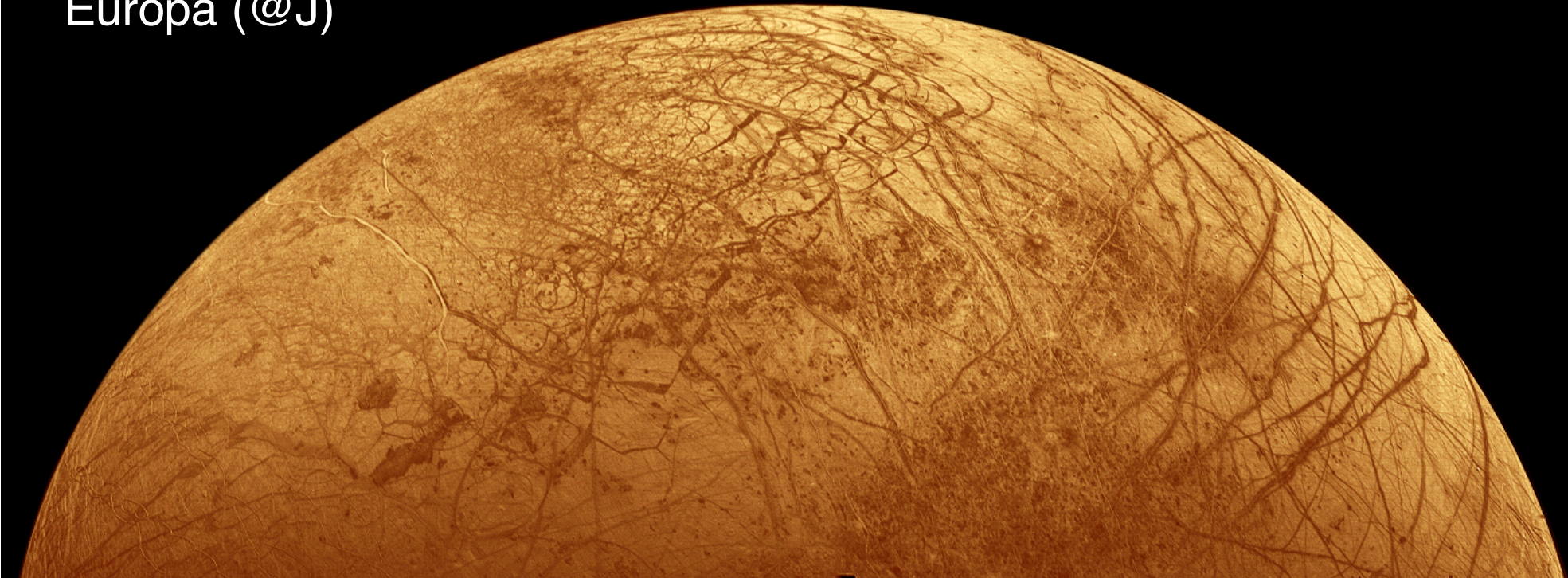
Rings full of waves (density)

Braided ring

Spokes

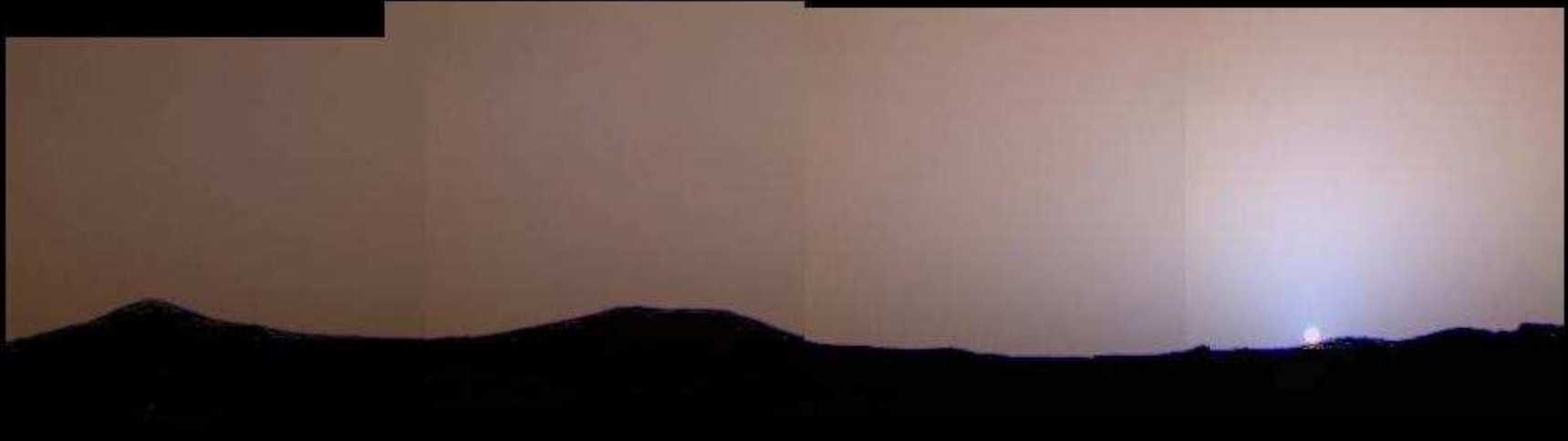


Europa (@J)



Planetary Atmospheres

- 1) Densities, temperatures
- 2) Origin of terrestrial planet atmospheres
- 3) Optics: colour, clouds
- 4) What happened to Venus?



Passively Heated by the Sun --- *the further the cooler*

Typically we observe objects in reflected light, however, all objects emit re-processed thermal radiation which is observable at longer wavelengths.

Blackbody temperature for a non-self-luminous spherical body at distance **a** away from the Sun (with albedo A -- reflectivity)

$$L_{\text{abs}} = (1-A) \frac{\pi R_p^2}{4\pi a^2} 4\pi R_s^2 \sigma T_s^4; \quad L_{\text{em}} = 4\pi R_p^2 \sigma T_p^4$$

If $L_{\text{abs}} = L_{\text{em}}$, then

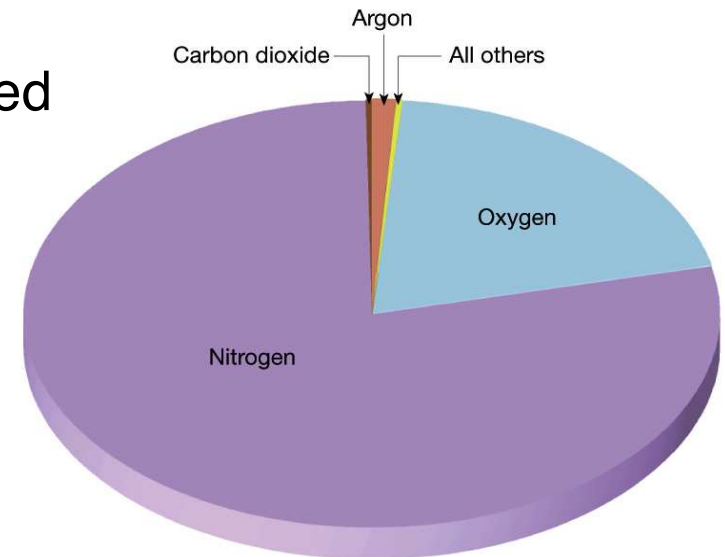
$$T_p = \left(\frac{R_o}{2a} \right)^{1/2} T_s (1-A)^{1/4}$$

	a (AU)	A	T _{pred} (K)	T _{act} (K)	
Mercury	0.4	0.06	422 K	100-725	(?)
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Uranus	19	0.51	60K	60	good
Neptune	30	0.62	40K	60	(?)
Comet at	5000	0.51	3.4K		

Atmospheres: Terrestrial Planets

	Atm. Composition	surface pressure/T	
Mercury	--	$< 10^{-12}$ bar	100-725 K
Venus	97% CO ₂ , 3% N ₂	92 bar	733 K (460°C)
Earth	78% N ₂ , 21% O ₂ , 1% Ar	1 bar	288 K (15°C)
Mars	95% CO ₂ , 3% N ₂ , 1.6% Ar	0.006 bar	223 K (-50°C)
Titan (@S)	95% N ₂ , few% CH ₄ , Ar	1.5 bar	93 K (-180°C)

Most atmospheres are reasonably well-mixed
(no molecular weight separation)



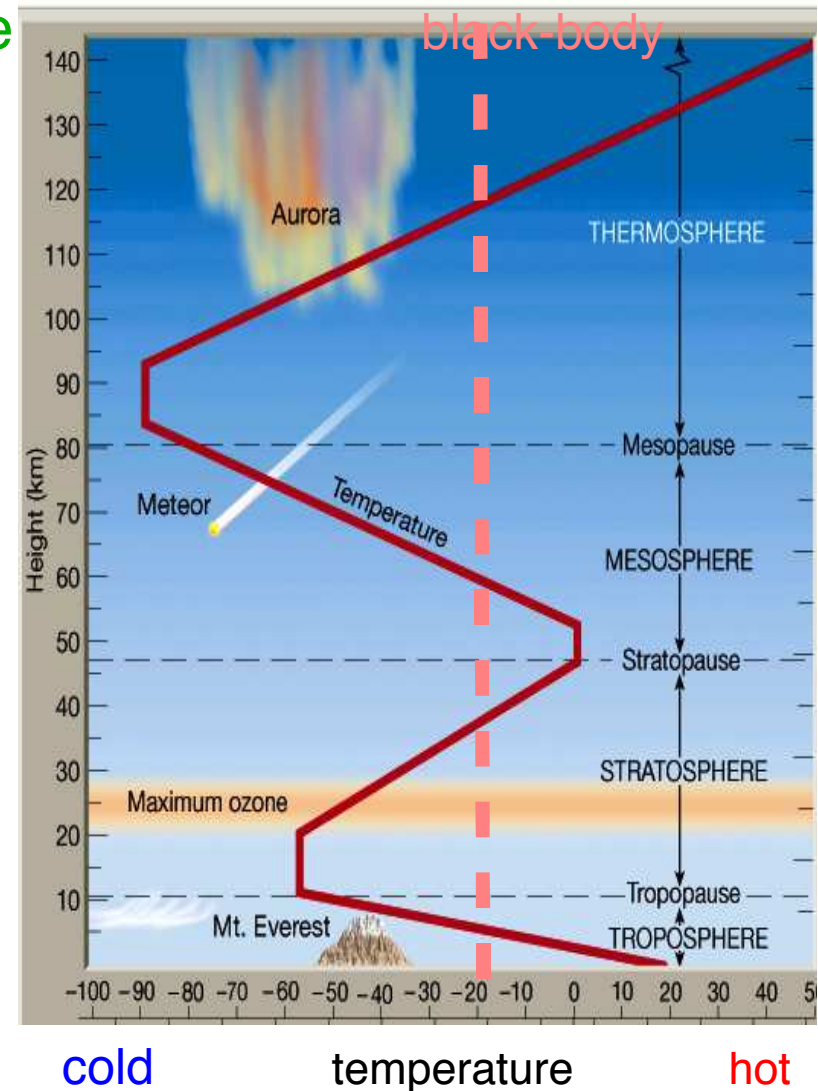
Earth's atmospheric composition

From http://www.ux1.eiu.edu/~cfjps/1400/atmos_origin.html

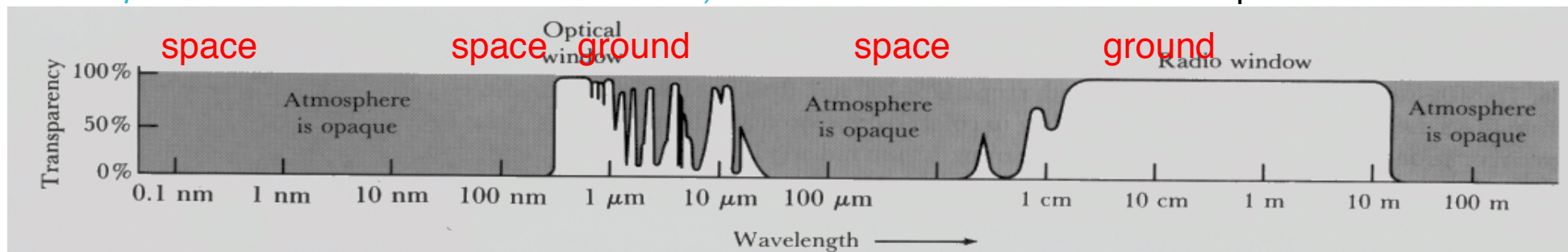
Density & Temperature of our atmosphere

- 1) Temperature largely **isothermal**;
density decreases exponentially, $H \sim 8$ km
Three local departures (T maxima)
 - Thermosphere absorbs X rays (~ 2000 K)
 - Stratosphere absorbs UV (O_3)
 - Ground absorbs whatever passes
- 2) Atmosphere largely transparent in optical,
but opaque in infrared --> green-house effect
 - Troposphere heated by ground--> turbulent
--> twinkling stars, planes fly @ ~ 10 km
 - Astronomical observations:
overcome turbulence & avoid absorption

(for Canadian Arctic site-testing, see
http://www.hia-ihc.nrc-cnrc.gc.ca/atrgv/inuksuit_e.html,
<http://www.casca.ca/ecass/issues/2006-ae/>)



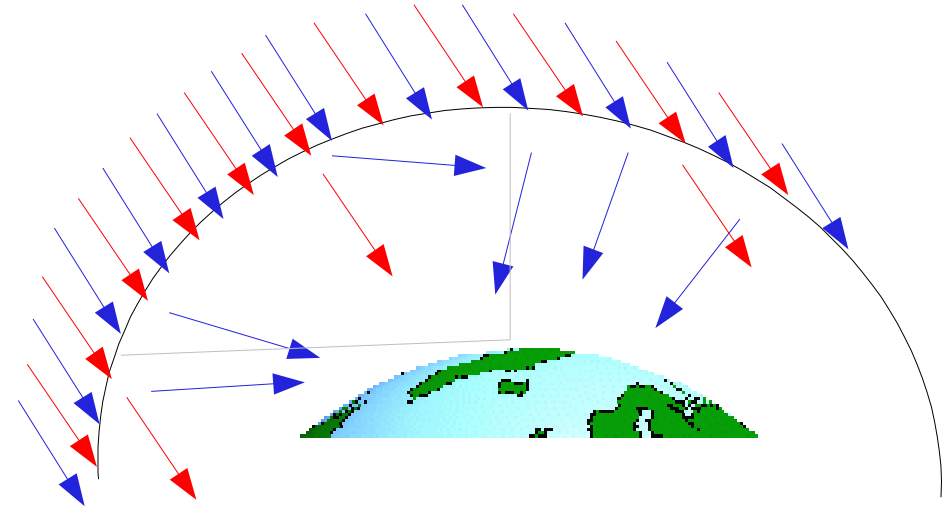
cold temperature hot



Atmospheric optics: 1) Why is the sky blue on Earth? Rayleigh scattering

air molecules & other constituents
(N₂, O₂, H₂O droplets, dust...) all have
sizes smaller than optical λ , and they
preferentially scatter short- λ photons:
 $\sigma \sim 1/\lambda^4$

Earth: *sky is blue (--> ocean blue)*
sunset is red (reddened)
horizon whiter than zenith
Fall/Winter sky dark blue
UV is diffuse



Moon: *sky is black*

Mars: *sky is reddish yellow*

fine-dust (1-10 μ m) Mie scattering --> white
iron oxide mineral absorption in the blue --> reddish

	x	y	sRGB pixel color
Sun above atmosphere	0.3259	0.3379	#fff3ea
5770 K blackbody (a Sun approximation)	0.3287	0.3397	#fff1e6
Illuminant B ("direct sunlight")	0.3840	0.3516	#ffbfaa

Mars Pathfinder true-color picture of Martian noon



Atmospheric optics: II) Clouds

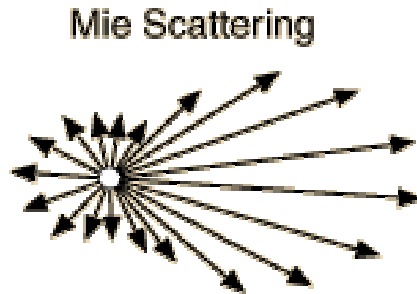
What are clouds?

Aggregates of water or ice droplets suspended in air
In troposphere: low clouds-- water; high clouds-- ice
100% hum. + condensation nuclei (dust, cosmic-rays)
e.g., rising air that cools (--> humidity increases)

How do they form?

Why are clouds white?

Water droplet colorless, solar light white
Mie scattering (droplets size $r \sim 10\mu\text{m} > \lambda$),
nearly geometric optics, no λ dependence
(at sunset, cloud is red)
soap foam: geometric scattering, also no λ dep.



Why don't clouds fall from the sky?

Tiny droplets, fall slowly; updraft mixing?

Fall and evaporate and form new ones?

Electrically charged clouds?



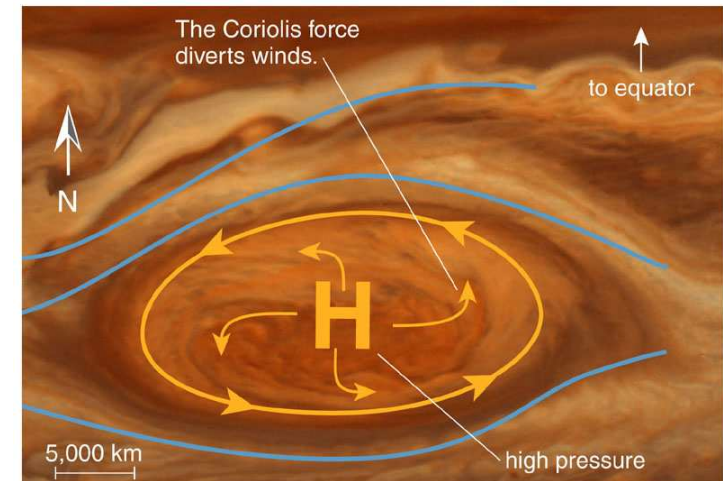
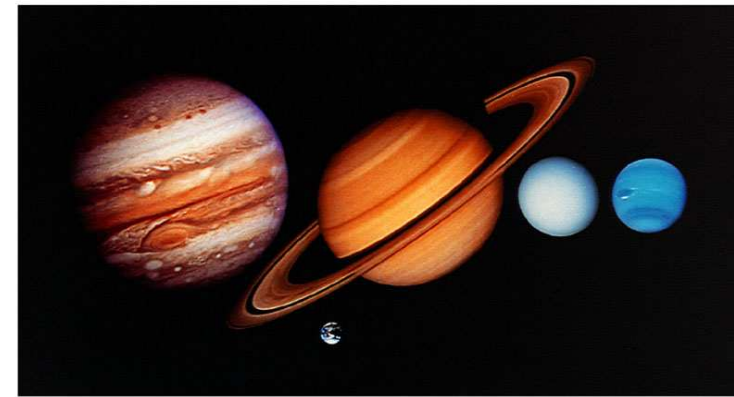
Intermezzo: Gas giant atmospheres

All 4 have deep atmospheres with mostly H₂ & He

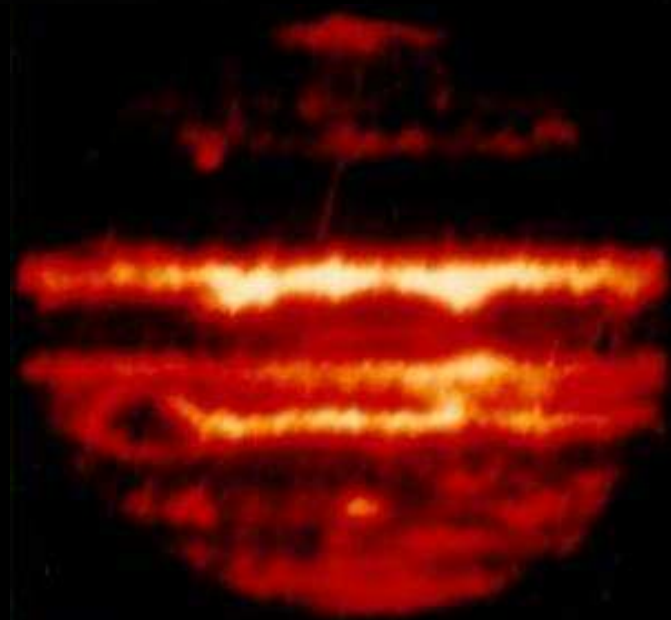
(fractions in % by volume, not by mass)

	J	S	U	N	Sun
H	88	97	83	74	86
He	11	3	15	25	14
CH ₄	0.2		2	1	
0.02 NH ₃					
0.0001 H ₂ O					

helium settling no helium settling



- Trace gases condense into **clouds** at diff. temperature
Clouds are also passive tracers of local wind pattern
- Jupiter, Saturn & Neptune have strong **zonal winds**
(up to 500 m/s)
zonal winds driven by solar irradiation,
a combination of cold pole-- hot equator pressure gradient & Coriolis force:
great red-spot of Jupiter: a giant anti-cyclonic vortex, surprisingly long-lived
cyclone: $2 \mathbf{V} \times \boldsymbol{\Omega} = - \text{grad } P/\rho$; tornado: $V^2/r = - \text{grad } P/\rho$
- Uranus: uniquely bland & sedate (no internal heat flux, obliquity 97 deg)



Origin of Earth's atmosphere

Our (& Venusian) atmosphere cannot be primordial

- 1) N_2 , CO_2 , H_2O are not condensed at 1AU from Sun,
 O_2 does not naturally occur
- 2) Earth too low in mass to accrete gas directly
- 3) Gas is unlikely to have been trapped in solids and dragged to Earth, since noble gases (Ne, Kr, Xe) are heavily depleted relative to solar abundance.
- 4) New-born Earth molten and hot (10^3K)
--> most gases can escape thermally.

Some relief only in that in the early bombardment period (~ 700 Myr) water can be brought in by comets & asteroids.

(Note: D/H ratio in comets ~ 2 higher than ocean, so these cannot do it alone)

Origin of Earth's atmosphere (cont'd)

Our atmosphere is obtained gradually: volcanic outgassing & invaders

1st atmosphere
thermal escape

H & He(?)

P: ?
T: $\sim 10^3\text{K}$

2nd atmosphere
outgassing/accretion

CO₂/NH₃ outgassed

H₂O accreted/outgassed

(solid crust/ocean, 3.5 Gyrs ago)

~ 100 bar (like Venus!)

$0^\circ\text{C} < T < 100^\circ\text{C}$

3rd atmosphere
absorbing CO₂

most H₂O liquid

CO₂ got locked in

O₂ produced

~ 1 bar

$\sim 15^\circ\text{C}$

sinks of CO₂: sedimentary rock via H₂O, life (carbon) via photon-synthesis

sources of CO₂: volcanic outgassing (+human activities)

sinks of H₂O: subducting plates

sources of H₂O: outgassing, comets/asteroids?

Currently sensitive balance reached, mild green-house

run-away green-house: too much CO₂, H₂O can all disappear -->

--> sink disappears as well while outgassing produces yet more CO₂

Venus: divergent evolution from Earth

	a(AU)	mass(M_E)	spin	atm. Pressure	T	tectonics	ocean
Earth:	1	1	1 day	1bar	288 K	Yes	Yes
Venus:	0.7	0.8	243 day	92bar	770 K	No	No

1) 97% CO₂ in the atmosphere, ~ 700K, *no CO₂ sink due to dryness*

2) Why so dry? high D/H ratio indicates past large H₂O reserve
Green-house runaway and H₂O photo-evaporated

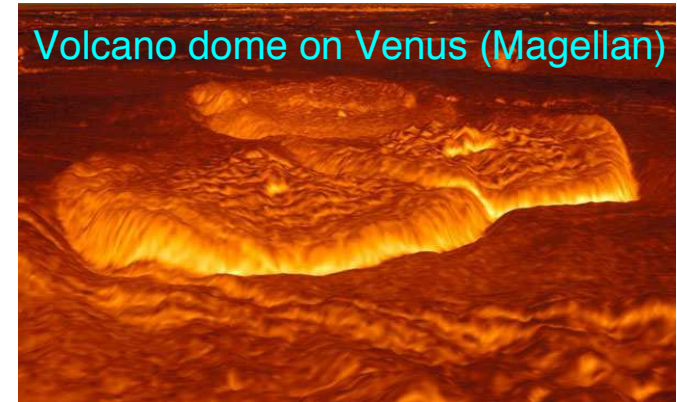
3) Cratering no older than ~0.8 Gyr --> tectonics stopped recently

A planet is a nonlinear system.
Strongly divergent evolution can occur.

Cause & Effect?

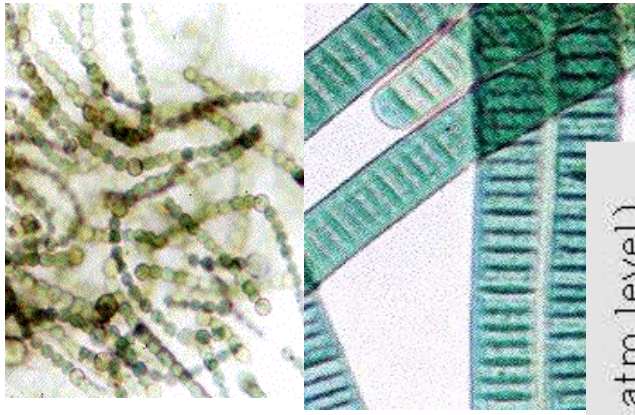
1) Slightly closer to the Sun and got torched?
Or formation site had naturally less H₂O?

2) Too much CO₂ to start with and H₂O never condensed
(But: Initial Earth atm. ~100 bar, mostly CO₂ --> would require *fine tuning*?)



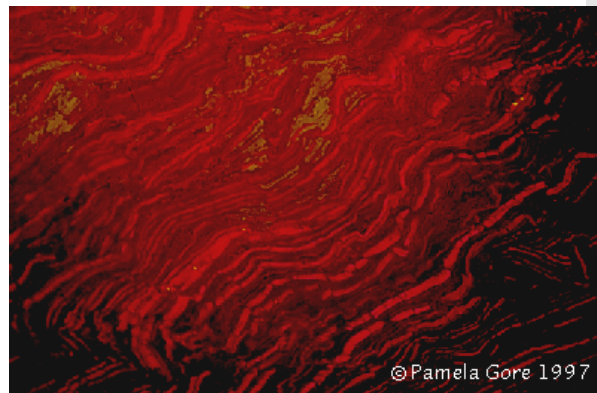
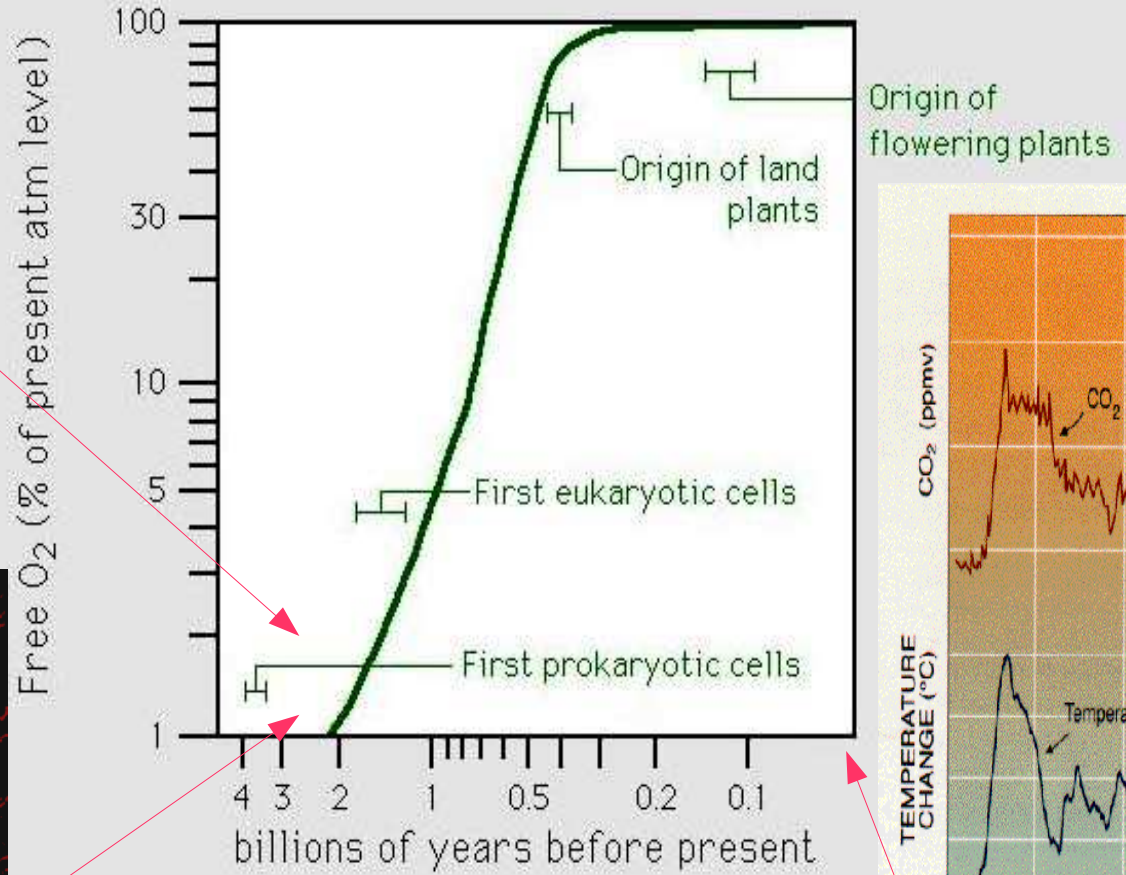
The Story for Mars: 2nd atmosphere gradually lost, no outgassing (tectonics)

Origin of O₂ on Earth: photosynthesis; $\text{CO}_2 + \text{H}_2\text{O} + h\nu \rightarrow \text{O}_2 + \text{carbo-hydrate}$



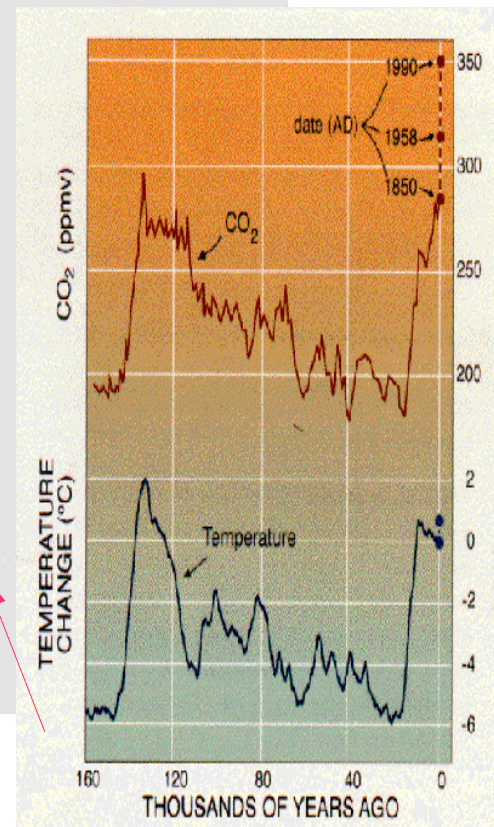
Archean Life: blue-green 'algae' or cyanobacteria (3.5-2.2 BYA) anerobic

From <http://www.clas.ufl.edu/users/mrosenme/Oceanography>



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Red-banded un-oxidized iron-rich rocks, pre-cambrian, ~2.5BYA
<http://www.dc.peachnet.edu/~pgore/geology/geo102/precamb.htm>



CO₂ and atm. T correlation
 (April 1989, Scientific. American)